



PALEOECOLOGICAL SIGNIFICANCE OF PALYNOMORPHS FROM PALEONTOLOGICALLY-DATED MAASTRICHTIAN–DANIAN DEPOSITS, BIR ABU MINQAR AREA, SOUTH WESTERN DESERT, EGYPT

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ABSTRACT

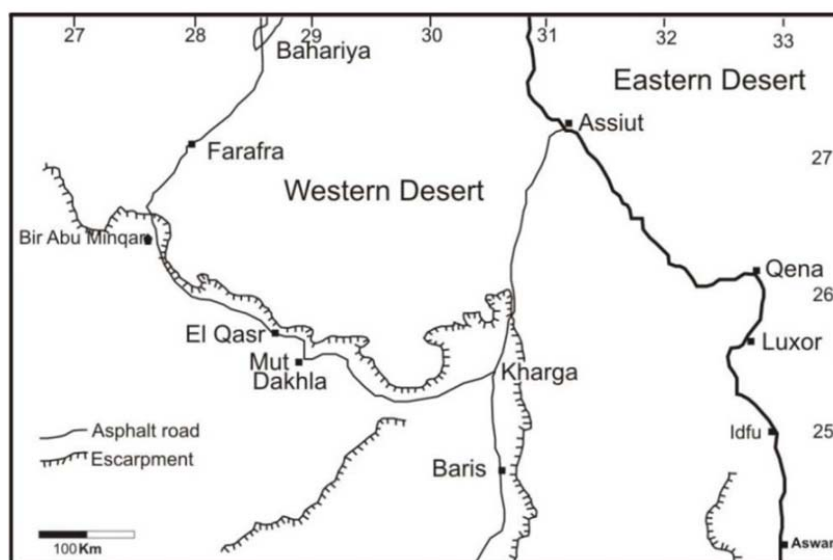
This study is carried out on the Dakhla Formation deposits at the Bir Abu Minqar section, South Western Desert, Egypt. Maastrichtian/Danian palynofloras have been recovered from the micropaleontology (foraminifera–nannoplankton)- and macropaleontology (bivalvia-ammonites)- dated deposits of the upper Dakhla Formation. Pteridophytic and liverworts spores, pollen of gymnosperms and angiosperms as well as dinoflagellate cysts have been identified. Angiosperm pollen includes typical members of the Late Cretaceous *Palmae* Province such as *Spinizonocolpites echinatus*. Presence of palm pollen and hygrophilic pteridophytic spores suggest a prevalence of humid tropical conditions during the deposition of the studied section, with *Nypa* palm mangroves thriving in the intertidal swamps that bordered the Late Cretaceous shallow sea at the area of the Abu Minqar.

Keywords: Maastrichtian–Danian, paleoecology, palynomorphs, South Western Desert, Egypt.

INTRODUCTION

The Bir Abu Minqar section (26 ° 31` 05" N, 27° 39` 43" E), approximately 260 km northwest of Mut Dakhla at the bend of the Dakhla-Farafra road, is located in the north of the Abu Minqar Village in the Western Desert (Fig. 1). The deposits of this section belong to the Dakhla Formation (introduced by Said, 1961) of the latest Cretaceous–earliest Paleogene (Maastrichtian–Danian) age.

Fig. 1. Sketched map showing the location of the Birr Abu Mina section (modified from Tantaway et al., 2001).



Geological background

The Dakhla formation is of special interest because of the rhythmic deposition of shales and glauconite-rich facies. Numerous studies have been made on the geology, biostratigraphy and sedimentology of the deposits exposed in the Dakhla Basin (e.g. Said, 1961, 1962; Hermina et al., 1961; Abbas and Habib, 1969; Issawi, 1972; El- Dawoody and Zidan, 1976; Omara et al., 1976, 1977; Barthel

and Hermann-Degen, 1981; Mansour et al., 1982; Faris, 1984; Hendriks et al., 1987; Luger and Schrank, 1987; Luger, 1988; Hermina, 1990; Kassab and Zakhera, 1995; Kassab et al., 1995; Tantawy et al., 2001). The deposits of the Bir Abu Minqar section are composed primarily of silty shale and thick calcareous sandstone beds that are rich in *Exogyra overwegi* (Bivalvia) and sphenodiscid ammonites (Barthel and Herrmann-Degen, 1981). *Exogyra overwegi* zone is approximately equivalent to the lower part of the planktonic foraminiferal zone CF7 and the lower part of the calcareous nannofossil zone CC25 (Tantawy et al., 2001). According to Tantawy et al. (2001, fig. 9), the Bir Abu Minqar section has the deposits of the Lower and lower Upper Maastrichtian, and after a hiatus the lower part of Danian. This is equivalent respectively to the planktonic foraminiferal zones CF7, CF6, and P1d (or P2), or the calcareous nannofossil zones CC 25a, CC25b, and NP4, spanning from about 70 to 69 Ma and 61-62Ma. This study is the first palynological study for the Bir Abu Minqar deposits. A future comprehensive set of samples from the studied deposits will be analyzed in a high-resolution work to infer possible biostratigraphic significance of palynomorphs. Since the current study is devoted essentially to contribute some paleoecologic significance of palynomorphs.

MATERIAL AND METHODS

In this study, 15 surface samples from Bir Abu Minqar area were studied. The deposits of the Dakhla Formation at the Bir Abu Minqar section are composed primarily of silty shale and thick calcareous sandstone beds (Fig. 2). 7 samples are productive only and yielded palynomorphs. They were prepared using the standard processing techniques (e.g. Wood et al., 1996). The semi-quantitative analysis of the recovered palynomorphs is presented in figure 2. Photomicrographs of selected spores, pollen and dinoflagellate cyst specimens have been compiled as Plate 1. The sample material and organic residues are housed in the collections of the Nanjing Institute of Geology and Palaeontology, Chinese Academy of Sciences, China. The microscope slides, primary data and figured material are housed in the collections of the Geology Department, Faculty of Science, Aswan Univ., and Aswan, Egypt.

RESULTS

Sparse and fairly preserved palynomorph associations (Plate 1) are recovered. Pollen and spores are the dominated groups (Fig. 2). The most common spore and pollen genera are *Deltoidospora* spp., *Triplanisporites* sp., *Inaperturopollenites* spp., *Spheripollenites* sp., *Araucariacites australlis*, and *Spinizocolpites* sp. The known ranges of selected palynomorphs recovered from the Bir Abu Minqar section (Fig. 2) suggest a Maastrichtian age for the palynologically productive parts of the studied section. Angiosperm pollen of stratigraphic significance include typical member of the Late Cretaceous Palmae Province (Herngreen et al., 1996) such as *Spinizocolpites* sp. (samples AM-4, AM-6 and AM-10, Fig. 2). In the North African region, several important Campanian–Maastrichtian angiosperm taxa were recorded such as *Spinizocolpites* (Schrank, 1984; El Beialy, 1995; Mahmoud, 2003; Mahmoud and Schrank, 2007). The triporite pollen is the most biostratigraphically significant genus recognized in this study. It has been reported from the Egyptian part of the Northern Gondwana Realm from the upper Cenomanian up to the Maastrichtian (Schrank, 1994; Schrank and Ibrahim, 1995; Ibrahim, 1996; Schrank and Mahmoud, 1998; Mahmoud and Schrank, 2007; Deaf et al., 2014).

Dinoflagellate cysts are poor; they are identified only from sample AM-14. The most common dinoflagellate cysts are *Palaeoperidinium cretaceum*, *Subtilisphaera senegalensis*, *Subtilisphaera* sp., and *Oligosphaeridium* sp. The dinoflagellate cyst taxa in this interval have relatively long range extending from the Early to the Late Cretaceous. *Subtilisphaera senegalensis* was originally described from the Aptian of Senegal (Jain and Millepied, 1973). It was later recoded from the Aptian-Albian of Egypt (e.g. Schank and Ibrahim, 1995; Ibrahim, 1996; Deaf et al., 2014). Similarly, it was frequently identified from Aptian rocks in Northwest Africa, specifically in Morocco and Libya (e.g.: Below, 1984; Thusu and Van der Eem, 1985). However, this species was found to range up to the Maastrichtian in the far eastern part of the Tethys (Shaozhi and Geoffrey, 1988). *Palaeoperidinium cretaceum* was also reported in Egypt to range from the Aptian to the Maastrichtian (e.g.: El Beialy, 1995; Deaf et al., 2016).

Paleoecological significance of Palynomorphs

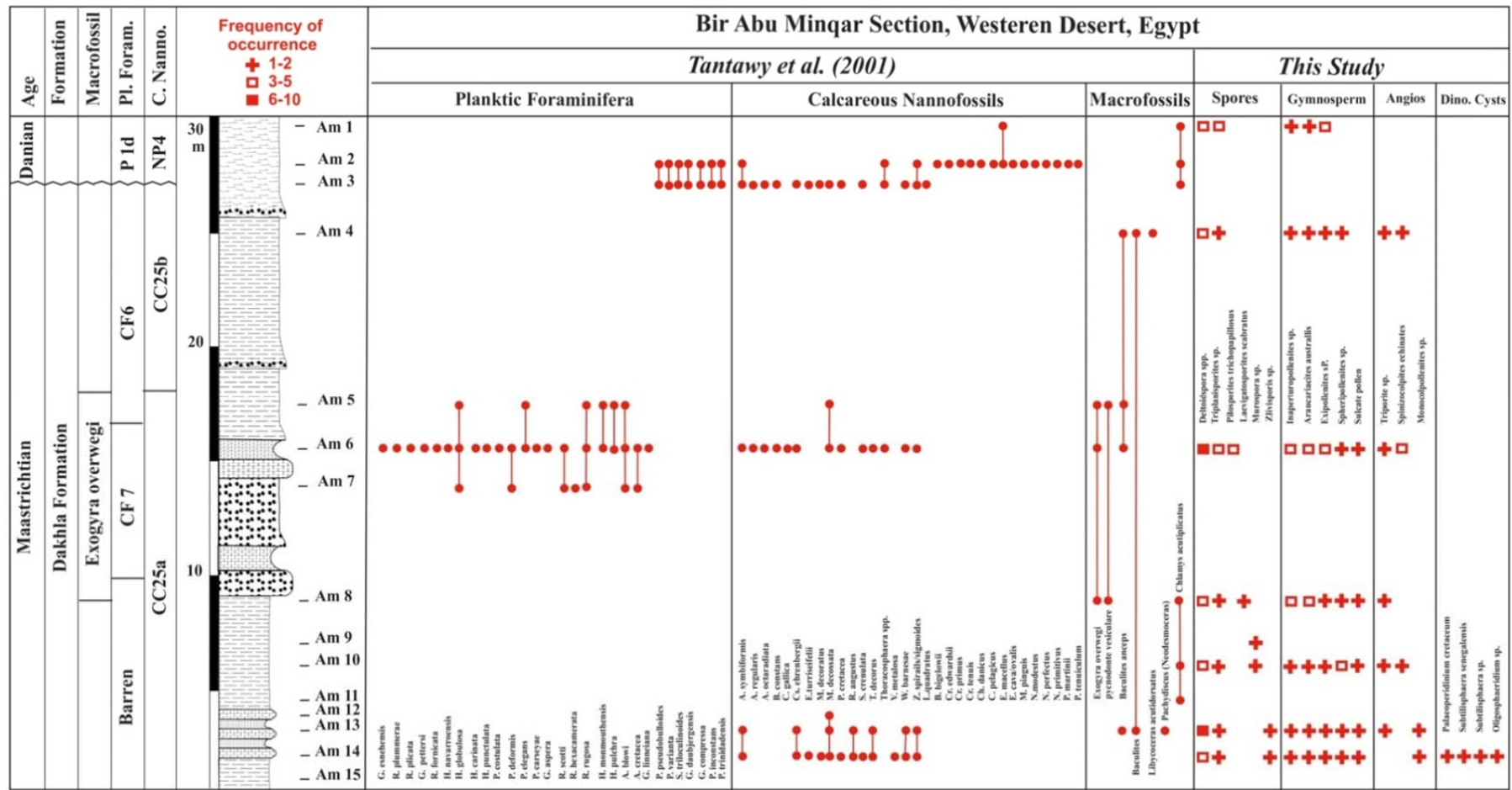


Fig. 2: Micropaleontology and macropaleontology age calibration of the Bir Abu Minqar section (after Tantway et al., 2011) and The semi-quantitative analysis of the recovered palynomorphs in the current study. (+ 1- 2 Rare, □ 3-5 present, ■ 6-10 common).

PALEOECOLOGICAL IMPLICATIONS

Paleoecologic and palaeoclimatic considerations in this section are based essentially on composition of palynoflora (Fig. 2). In most of the studied samples, the present to common abundances of the hygrophilous palynomorphs (mainly fern spores) such as *Deltoidospra* and rare to present abundances of the (liverworts spores) *Zlivisporis* probably reflect the local pteridophyte vegetation and wetlands (Schrank and Mahmoud, 1998; Mahmoud and Schrank, 2007; Deaf et al., 2014, 2020). In addition, the rare to present abundances of the araucariacean pollen (*Araucariacites australlis*) reflect conifer vegetation on relatively dry hinterlands (Schrank and Mahmoud, 1998; Mahmoud and Moawad, 2002; Deaf et al., 2020).

Pollen of palms is an important element of the terrestrial palynoflora of the Bir Abu Minqar section. Palms may have been abundant in tropical humid lowlands of the coastal plain where they were associated with pteridophytes and other plants that inhabited lowland sites (e.g. Schrank and Mahmoud, 1998; Deaf et al., 2014, 2020). The presence of a *Nypa* mangrove belt is indicated by the rare to present abundances of the *Spinizonocolpites* (Fig. 2). In Egypt, this is the southernmost occurrence of this widespread mangrove type, which was previously recorded from northern and central Egypt (e.g. Schrank, 1987; Schrank and Ibrahim, 1995; Mahmoud and Schrank, 2007; El-Soughier, 2011).

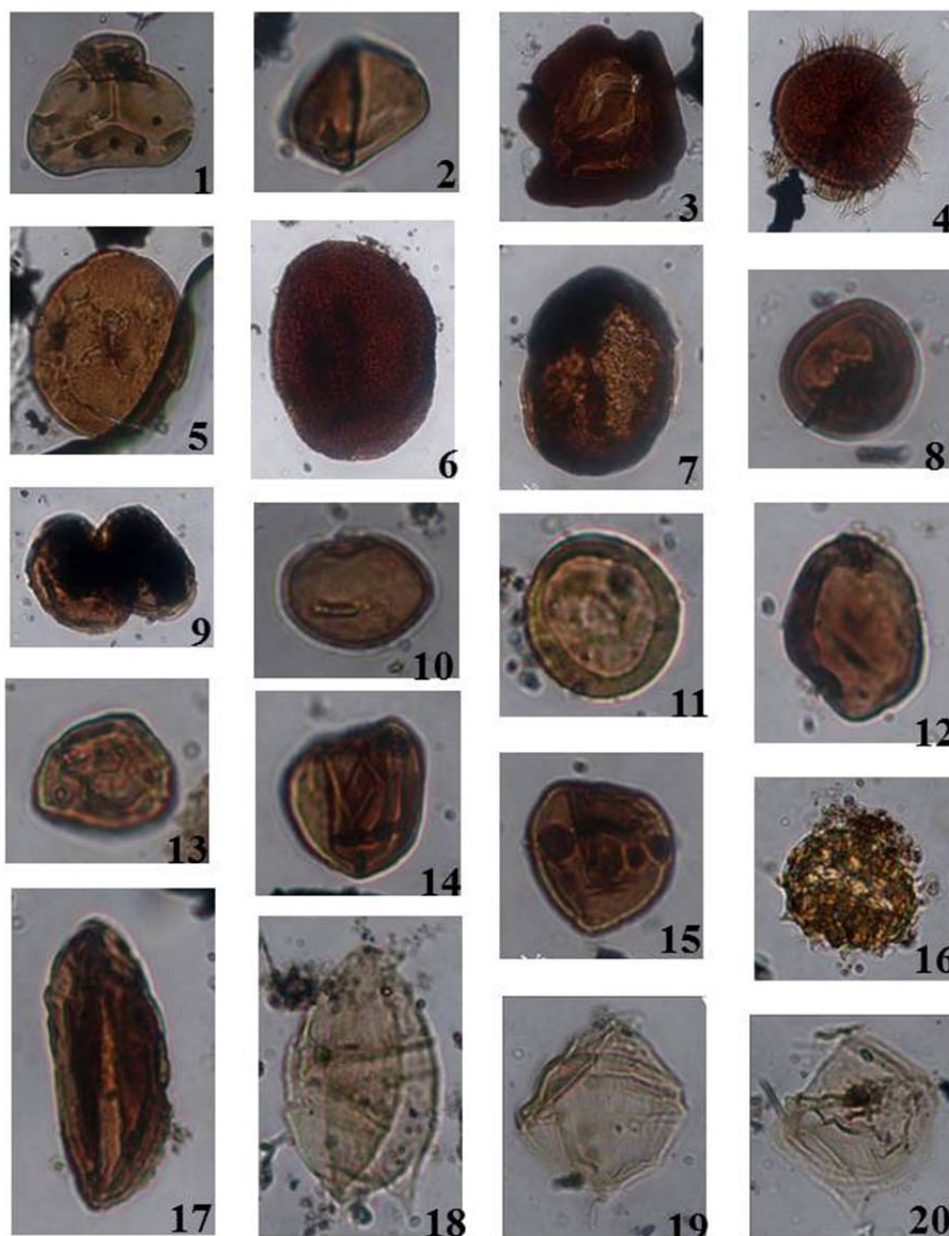
It is important to mention that the Late Cretaceous period was characterized by a relatively warmer climate and high eustatic sea levels (Stanley, 1999) in contrast to the early Paleocene that was relatively cooler and drier (Gould, 1993). *Nypa* vegetation reached its peak during the Eocene when the climate shifted towards warmer conditions (Corner, 1966). *Nypa* mangrove was recorded to dominate the swamp settings in West Africa during the Eocene (Sowunmi, 1986). By the end of the Eocene, the climate shifted towards cool conditions, which became unfavorable for *Nypa* to survive and eventually vanished. Based on the data mentioned above, the occurrence of the *Nypa* mangrove pollen (*Spinizonocolpites*) suggests the development of warm humid swamp conditions during deposition of the studied Upper Cretaceous rocks. The mostly rare abundances of the hot dry indicator pollen grains *Exesipollenites* sp. and *Inaperturopollenites* spp. also support the interpretation of the prevalence of regional humid conditions at the study area (e.g. Doyle et al., 1982; Schrank and Mahmoud, 1998; Deaf et al., 2020).

In the lower part of the section, the rare abundances of the peridinioid dinoflagellate cysts *Palaeoperidinium* and *Subtilisphaera* reflect marginal marine environment (e.g. Mahmoud and Schrank, 2007; Deaf et al., 2014, 2020). This interpretation is consistent with conclusion derived from sedimentological, mineralogical, and micro- and macrofossil evidences from the Dakhla Basin including our studied area (Fig. 2). During the Maastrichtian-early Paleocene, a shallow sea inundated the Western Desert of Egypt and reached as far south as of the Dakhla Basin (Tantawy et al., 2001). Tantawy et al. (Op. Cit) suggested that regional climate shifted towards more humid conditions during the latest Maastrichtian–earliest Paleocene as is evidenced from the increased kaolinite deposition. Thus, it is suggested that the *Nypa* palm mangroves recoded herein most probably grew in the intertidal swamps that bordered this shallow sea at the area of Abu Minqar.

CONCLUSION

1. 18 species of Maastrichtian-Danian palynofloras were recovered from the Bir Abu Minqar section. The productive samples favor a Maastrichtian age as is established by macro- and microfossils.
2. The most common palynofloras are *Deltoidospora* spp., *Triplanisporites* sp., *Inaperturopollenites* spp., *Spheripollenites* sp., *Araucariacites australlis*, and *Spinizocolpites*.
3. Presence of palm pollen along with the hygrophilic pteridophytic spores related to water ferns and hepatics suggests a prevalence of humid tropical conditions during the deposition of the upper Dakhla Formation deposits at the Bir Abu Minqar area. *Nypa* palm mangroves probably grew in the intertidal swamps bordering the Late Cretaceous (Maastrichtian) shallow sea at the area of Abu Minqar.

Plate



Magnifications are 600X (1-11, 17-20); 1000X (12-16)

50 μ m

Plate. 1: Spores (1-4), pollen (5-17) and dinoflagellate cysts (18-20) of Danian-Maastrichtian Bir Abu Minqar section (AM). Magnifications are 600X (1-11, 17-20); 1000X (12-16).

1. *Deltoidspora* sp., AM- 4, Diameter 62 μ m. 2. *Triplanisporites* sp., AM-1, Diameter 56 μ m. 3. *Murospora* sp., AM-10, Diameter 82 μ m. 4. *Pilosporites trichopapillosus* (Thiergart) Delcourt & Sprumont, 1955, AM-6, Diameter 68 μ m. 5, 8, 11, 15. *Inaperturopollenites* sp., AM- 8, Diameter 40 μ m, AM-13, Diameter 88 μ m. 6. *Punctatosporites scabratus* (Couper) Norris, 1965, AM- 8, Diameter 94 μ m. 7. *Araucariacites australlis* Cookson ex Couper, 1953, AM-14, Diameter 92 μ m. 9. *Spheripollenites* sp., AM-14, Diameter 32 μ m. 10. *Expipollenites* sp., AM-13, Diameter 34 μ m. 12, 14. *Proteacidites* sp., AM- 13, Diameter 32-44 μ m. 16. *Spinizocolpites echinatus* Muller, 1968, AM-10, Diameter 36 μ m. 17. *Monocolpopollenites* sp., AM-13, Diameter 90 μ m. 13. *Zlivisporis* sp. sensu Mahmoud and Schrank, 2007, 18. *Subtilisphaera* sp., AM-14, Length 90 μ m. 19. *Subtilisphaera senegalensis* Jain & Millepied, 1973, AM-14, Length 58 μ m. 20. *Palaeoperidinium cretaceum* Poccok, 1962, AM-14, Length 67 μ m, breadth 68 μ m.

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Appendix

List of species

Spores and pollen grains

Araucariacites australlis Cookson ex Couper, 1953

Deltoidspora spp.

Exesipollenites sp.

Inaperturopollenites spp.

Punctatosporites scabratus (Couper) Norris, 1965

Monocolpopollenites sp.

Murospora sp.

Pilososporites trichopapillosus (Thiergart) Delcourt & Sprumont, 1955

Triplanisporites sp.

Proteacidites sp.

Spinizocolpites echinatus Muller, 1968

Zlivisporis sp.

Dinoflagellate cysts

Palaeoperidinium cretaceum Poccok, 1962

Oligosphaeridium sp.

Subtilisphaera senegalensis Jain & Millepied, 1973

Subtilisphaera sp.

Paleoecological significance of Palynomorphs

الأثار البيئية للحفريات النباتية من رواسب المسترختي-الدانيان المعرفة بالحفريات-منطقة أبو منقار- جنوب الصحراء الغربية -مصر

ماهر ابراهيم الصغير

الخلاصة

قسم الجيولوجيا- كلية العلوم - جامعة اسوان

أجريت هذه الدراسة على رواسب تكوين الداخلة في منطقة بئر أبو منقار ، جنوب الصحراء الغربية ، مصر. تم التعرف على حفريات النباتية لرواسب Maastrichtian/Danian المعرفة اعتمادنا على الحفريات الدقيقة (foraminifer nanoplankton) -وحفريات الكبيرة (bivalvia – ammonites) الرواسب المؤرخة في تكوين الداخلة العلوي . تم التعرف على الأبواغ والنباتات المحبة للماء، وحبوب اللقاح والسوطيات. يشمل لقاح Angiosperm على Palmae الموجودة في المتأخر الطباشيري مثل Spinizonocolpites echinatus. يشير وجود حبوب اللقاح وأنواع الأبواغ انتشار الظروف المدارية الرطبة أثناء ترسب القسم المدروس ، مع ازدهار أشجار نخيل Nypa في مستنقعات المد التي تحد البحر الضحلة أواخر العصر الطباشيري في منطقة أبو منقار.